

1                    **Supplementary Information for:**

2                    **Plate motion drives variability in ocean oxygenation**

3                    **through the Phanerozoic**

4

5                    **Alexandre POHL<sup>1,2\*</sup>, Andy RIDGWELL<sup>1\*</sup>, Richard G. STOCKEY<sup>3</sup>, Christophe**  
6                    **THOMAZO<sup>2,4</sup>, Andrew KEANE<sup>5</sup>, Emmanuelle VENNIN<sup>2</sup>, Christopher R. SCOTSESE<sup>6</sup>**

7                    *<sup>1</sup>Department of Earth and Planetary Sciences, University of California, Riverside, CA,  
8 USA*

9                    *<sup>2</sup>Biogéosciences, UMR 6282, UBFC/CNRS, Université Bourgogne Franche-Comté, 6  
10 boulevard Gabriel, F-21000 Dijon, France*

11                    *<sup>3</sup>Department of Geological Sciences, Stanford University, Stanford, CA 94305, USA*

12                    *<sup>4</sup>Institut Universitaire de France, Paris, France*

13                    *<sup>5</sup>School of Mathematical Sciences and Environmental Research Institute, University  
14 College Cork, College Road, Cork, Ireland*

15                    *<sup>6</sup>Department of Earth and Planetary Sciences, Northwestern University, Evanston,  
16 Illinois 60208, USA*

17

18                    *\*These authors contributed equally to this work*

19

20

21

22 **Supplementary Discussion**

23 Most of the simulations in Extended Data Figs. 5 and 6 show that the ocean dynamics  
24 eventually settles to a stable steady-state solution. However, amongst these are cases where  
25 the system overshoots (for example, compare panels a and g of Extended Data Fig. 5) before  
26 approaching a steady-state solution. This implies the presence of an oscillatory mode, in  
27 agreement with the linear stability analysis of low-resolution GCM steady states under  
28 present-day conditions<sup>47,48</sup>. The simulations in Extended Data Figs. 5 and 6 reveal that,  
29 depending on changes to continental configuration, the steady-state solutions can lose  
30 stability, resulting in stable oscillations. It is known that, under certain climate conditions,  
31 advective feedbacks in the ocean circulation can destabilize the steady-state<sup>49</sup>. Extended Data  
32 Figs. 5bb and 6z are cases of steady-state solutions with oscillatory modes that are linearly  
33 stable, but only weakly stable, resulting in clear examples of damped oscillations. Since the  
34 frequencies of the damped and self-sustained oscillations of Extended Data Figs. 6z and 6aa,  
35 respectively, are very similar, the transition could represent a Hopf bifurcation, resulting in  
36 self-sustained oscillations around an unstable steady state.

37 The precise conditions that allow not only the destabilization of the steady state, but also  
38 the existence of *stable* oscillations, as well as the precise bifurcation structure associated with  
39 the oscillations, is the subject of ongoing work with idealized continental configurations.

40

41 **Supplementary References**

42 47. Weijer, W. & Dijkstra, H. A. Multiple oscillatory modes of the global ocean  
43 circulation. *J. Phys. Oceanogr.* **33**, 2197–2213 (2003).

44 48. Sirkes, Z. & Tziperman, E. Identifying a damped oscillatory thermohaline mode in a  
45 general circulation model using an adjoint model. *J. Phys. Oceanogr.* **31**, 2297–2306  
46 (2001).

47 49. Weijer, W. *et al.* Stability of the Atlantic Meridional Overturning Circulation: A  
48 Review and Synthesis. *J. Geophys. Res. Ocean.* **124**, 5336–5375 (2019).

49