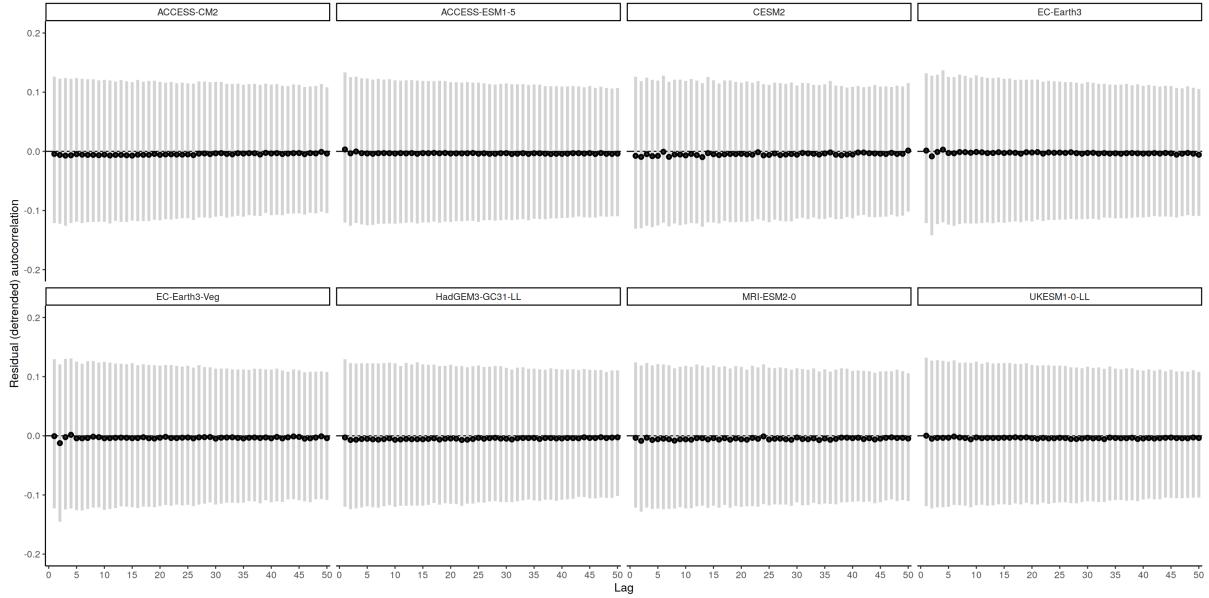


1321 **Supplementary Information**

1322 **S1 Conditions for validity of method**



SI Figure S1: Temporal autocorrelation of detrended residuals of Rx1d for lags from 1 to 50, for all eight CMIP6 models used. Detrending was done by subtracting the loess-filtered ensemble mean timeseries per gridcell.

1323 As mentioned in the main text, the temporal independence of Rx1d values is a prerequisite for
1324 temporal independence of record breaking probabilities and validity of equations (1), (2) and (3). Fig.
1325 S1 shows that autocorrelations between Rx1d values at both high and low frequencies (short and long
1326 lags) are 0. The autocorrelation of the non-detrended values is positive and gradually decaying because
1327 there is a long term trend. This trend affects the record breaking probabilities over time, but does not
1328 violate temporal independence of Rx1d values.

1329 **S2 Spatial pooling methods for observational/reanalysis GEV
1330 fits**

1331 As mentioned in Sect. 4.2, we tested two different spatial pooling methods to improve the GEV fits to
1332 the short observational and reanalysis timeseries. Below we outline the effects of shape-only and naive
1333 spatial pooling.

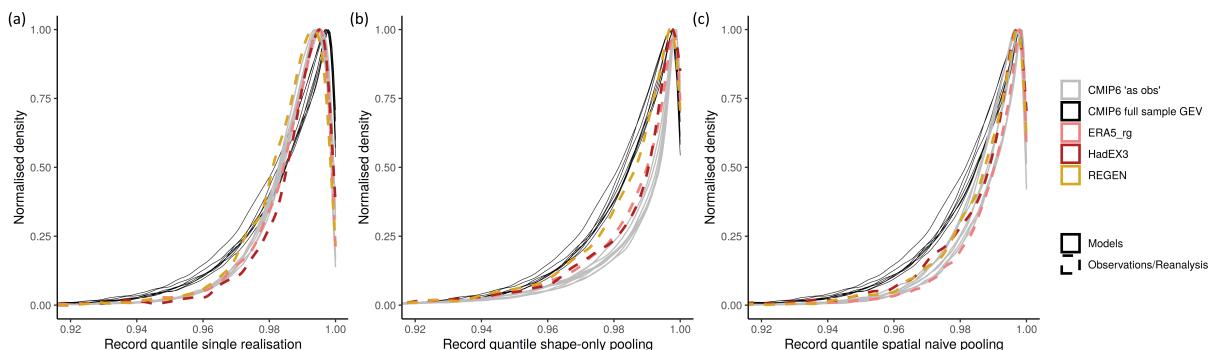
1334 For shape-only spatial pooling, strength is borrowed from neighbouring gridcells to improve the
1335 estimate of ξ only [88]. In practice, we pool Rx1d data within a spatial window of 5×5 gridcells to fit a
1336 GEV, determining unique μ and σ values for each gridcell in the window, but allowing only one universal
1337 ξ value for all gridcells in the window. This is achieved by defining spatial covariates for μ and σ that,
1338 as it were, ‘turn on’ the individual gridcells in the fitting procedure, while ξ has no spatial covariate.
1339 The resulting GEV parameters are assigned to the middle gridcell of the window. Naive spatial pooling,
1340 on the other hand, implies we simply fit one GEV to all the Rx1d data pooled in a spatial window, and
1341 assign the GEV parameters to the middle gridcell, as used by e.g. [89]. There is thus a degree of spatial
1342 smoothing for all three GEV parameters in this case, whereas the shape-only method aims at smoothing
1343 only ξ . [36] compares several GEV estimation methods including the two we test and finds that both
1344 are equally effective at improving accuracy of Rx1d return level estimates for return periods longer than
1345 the sample size.

1346 Supplementary Fig. S2 shows the effect of the pooling schemes on the distribution of quantiles.
1347 Supplementary Fig. S2a-c show the normalised density of record quantiles, where the black lines show
1348 the model means determined based on full-ensemble GEV distributions, i.e. the “target”, and grey lines

1349 show the model results when members are treated as HadEX3 observations. The coloured dashed lines
 1350 correspond to record quantiles in the observational and reanalysis data. Supplementary Fig. S2a shows
 1351 the quantiles obtained for GEV distributions fitted ‘as observations’ to timeseries without spatial pooling.
 1352 The peak lies too low, and also low quantiles are underestimated: the record quantile distribution is too
 1353 narrow. Below we discuss the effects of pooling and how the impact the record quantile distribution.

1354 Shape-only pooling leads to smoother patterns and smaller magnitudes of ξ , which makes them more
 1355 similar to the full-ensemble distribution’s ξ , see Supplementary Fig. S3b. However, also σ responds
 1356 strongly to this pooling scheme, since the setup aims to fit a σ and μ that are specific to single gridcell
 1357 data. One might say that shape-only pooling partly decouples the fit of μ and σ from the fit of ξ , where
 1358 the μ and σ are fit with only part of the information (one gridcell). This leads to compensation effects
 1359 in primarily the σ values, which strongly affect the estimated record quantiles. In part of the gridcells,
 1360 the compensation leads to decreases in σ relative to the single-gridcell fits, which are associated with
 1361 strong increases in quantiles estimated – the smaller the initial quantile, the stronger the increase. In
 1362 another subset of gridcells, the compensation leads to increases in σ , which leads to decreasing quantile
 1363 values, especially for already lower quantiles. The combination of these two σ -related quantile changes
 1364 leads to clustering of most of the quantiles at very high values, and some being moved to very low values.
 1365 Therefore we see a certain “pulling apart” in the quantile distribution which leaves a gap in the middle
 1366 ranges where quantiles between 0.95 and 0.99 should be found, Supplementary Fig. S2b. This effect is
 1367 seen for the observational and reanalysis record quantiles as well. Most of the quantiles are overestimated
 1368 in this approach, leading to a strong underestimation of record breaking probabilities. We tried different
 1369 pooling window sizes of 3×3 and 5×5 on the HadEX3 grid, and see that the 5×5 window size results
 1370 in smoother shape parameters, more similar to those of the full-ensemble fit, however, the discrepancy
 1371 in the quantile distributions increases with increasing window size.

1372 In the naive pooling setting, a coherent scale and shape is fit to the pooled sample, leading to less
 1373 local compensation effects of the scale parameter: we see much smaller changes relative to the single-
 1374 gridcell fit, whereas the smoothing of the shape parameter is almost as effective as in the shape-only
 1375 pooling setting. From Supplementary Fig. S2a to c we see improvement in the location of the peak and
 1376 and higher quantiles brought about by spatial pooling, and minor improvement for the lower quantiles
 1377 as well. Between quantiles 0.98 and 0.94 the offset remains however considerable, explaining the biases
 1378 in the CCP values, shown below in the validation section. Also for the observations and reanalysis
 1379 (coloured dashed lines) we see improvement in the peak location and a shape more similar to the black
 1380 full-ensemble target lines. The comparison here is imperfect since we do not know the true quantile
 1381 distribution of the observational and reanalysis record quantiles – the distribution can differ from the
 1382 model distributions due to model and observational errors, and differences in coverage.



SI Figure S2: Normalised density plots of the historical record quantiles determined using different GEV fitting methods. In all plots, solid grey lines show the ‘true’ record quantiles determined from full-ensemble GEV fits, and solid black lines show the model-as-observations comparisons, where GEV distributions were fitted ‘as observations’ to the simulation data. Dashed coloured lines correspond to the observational and reanalysis datasets. (a) shows the results for GEV distributions fitted to single gridcell data, (b) shows the results for GEV distributions fitted using shape-only spatial pooling, and (c) shows the results for GEV distributions fitted using naive spatial pooling. See also Sect. 4.2.

1383 We perform a few additional tests to confirm the seemingly better performance of naive spatial
 1384 pooling. As the aim of our study is to estimate record breaking probabilities, we assess the skill of
 1385 the probability estimate for each spatial pooling method using ranked probability skill scores (RPSS).
 1386 These are determined by estimating cumulative record breaking probabilities treating model members as

1387 observations, and comparing the estimate to the actual future evolution of the model member; see Sect.
 1388 4.4 on validation for a full explanation. The RPSS represents the improvement of the estimation method
 1389 in question relative to a benchmark. Table 3 in Sect. 4.4 shows RPSS values of the record breaking
 1390 probability estimates corresponding to different GEV-methods used to determine the record quantiles.
 1391 The first row refers to GEV distributions fitted to the 1950–2015 timeseries of single gridcells, where no
 1392 optimisation has been done to reduce biases in the GEV fit due to the small sample size. We see a minor
 1393 skill improvement of 5% over the benchmark. For shape-only spatial pooling, the improvement over
 1394 the benchmark is in fact negative, i.e. the probability estimates are worse. We tested different spatial
 1395 window sizes for the shape-only pooling, and see that 5×5 on the HadEX3 grid is the minimum size
 1396 to achieve clearly smoother patterns of ξ . For larger windows, the results are similar but the prediction
 1397 gets slightly worse as window size is increased. GEV fits using naive spatial pooling lead to more than
 1398 twice as much skill gain as the single gridcell GEV. We tested larger window sizes and see that a window
 1399 size of 3×3 on the HadEX3 grid is better than larger windows. The probability estimates based on full
 1400 ensemble GEV distributions are 30% more accurate than the benchmark.

1401 Lastly, we confirm that not just bulk properties, but also the spatial pattern of ξ and the probability
 1402 metrics improve most when naive spatial pooling is used. Supplementary Fig. S3 shows the ξ (a-d), record
 1403 quantile (e-g), state likelihood (i-l) and CCP (m-p) maps of one single member of ACCESS-ESM1-5 –
 1404 the largest ensemble in our model selection. We show metrics based on single gridcell, shape-only pooled,
 1405 naively pooled and full-ensemble GEV fits for visual comparison, and Supplementary table S4 provides
 1406 the multi-model spatial correlations of these quantities obtained from the different ‘as observations’ GEV
 1407 fits to those obtained using the full-ensemble GEV.

1408 The maps and spatial correlations show a much stronger agreement of the naive pooling GEV based
 1409 results with the full-ensemble GEV based results. For shape-only pooling, strong patterning appears
 1410 that seems influenced by the climatology and leads to artificial regions of low state likelihood and high
 1411 future record breaking probability.

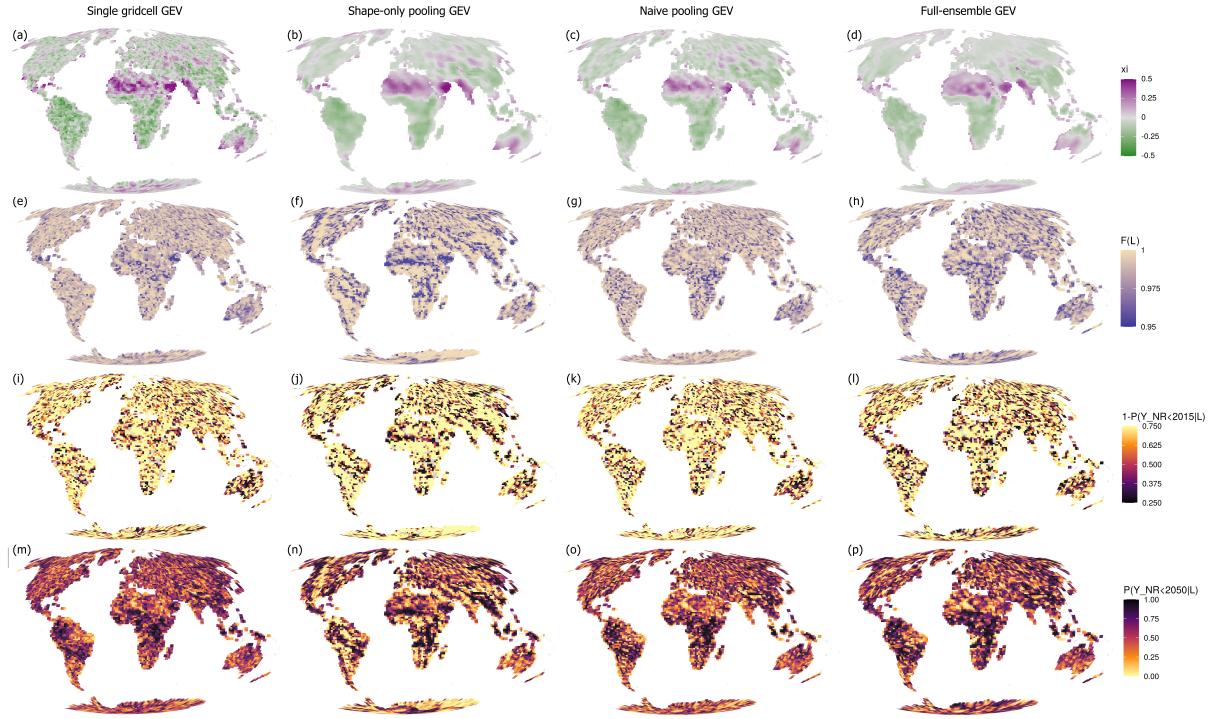
1412 Given the clearly better performance (for our purposes) of naive spatial pooling in all tests performed,
 1413 we employ naive spatial pooling for the observational/reanalysis GEV fits in our analysis.

SI Table S4: Spatial correlation coefficients of the variables listed in the top row; correlations of the result obtained using the GEV fitting method listed in the first column with the result obtained using the full-ensemble GEV. SL refers to state likelihood: $1 - P(T < Y_{NR} \leq 2015 | L)$ and CCP to $P(Y_{NR} \leq 2050 | L, T = 2015)$

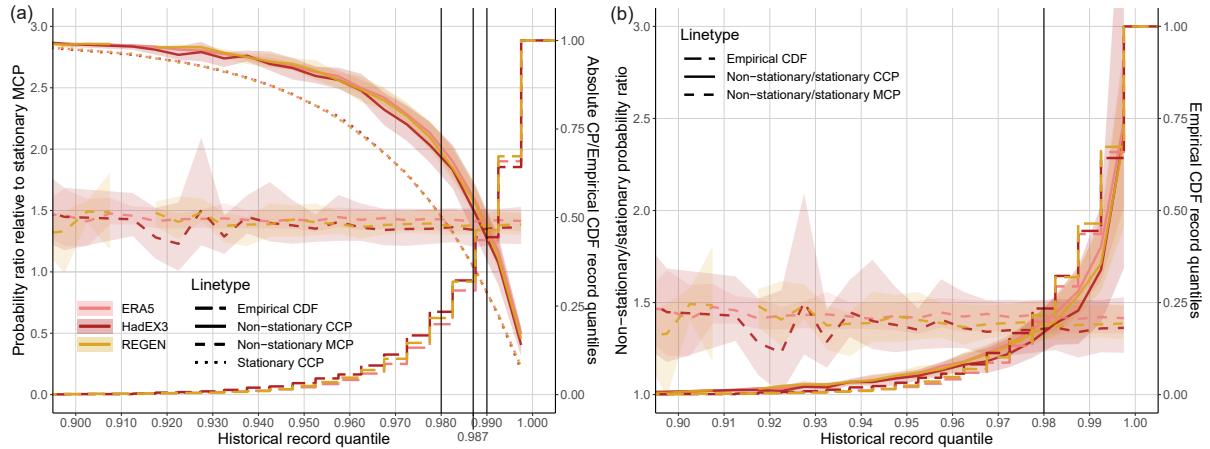
GEV fitting method	ξ	μ	σ	$F(L)$	SL	CCP
Single gridcell ‘as obs’ GEV	0.70	0.91	0.97	0.45	0.69	0.61
Shape-only spatial pooling GEV, 5×5 window	0.78	0.87	0.91	0.53	0.67	0.60
Naive spatial pooling GEV, 3×3 window	0.74	0.94	0.98	0.75	0.84	0.80

1414 S3 Climate model results - comparison of GEV fitting methods

1415 In Fig. S3 the sensitivity to the GEV fitting method is shown. In the first three columns the single
 1416 model member is treated as an observational record. The single gridcell GEV leads to too much spatial
 1417 variability, yet too little spread in the quantiles. The strong increase in spread but artificial pattern
 1418 resulting from shape-only pooling (see Sect. 4) is very apparent in the second column. The moderate
 1419 changes due to naive spatial pooling (third column) clearly do not reproduce the ‘true’ full-ensemble
 1420 patterns (right column), but are closest in terms of magnitude, spread, and spatial pattern. The full-
 1421 ensemble patterns, showing more distinct (random) regions of low record quantiles and state likelihood,
 1422 and high future record breaking probabilities are indicative of what any true observational pattern could
 1423 look like if we were able to determine the true underlying distribution. The pattern of record quantiles
 1424 is random, and elevated or damped record breaking probabilities due to natural variability can occur
 1425 anywhere.



SI Figure S3: Patterns of GEV shape parameter ξ (a-d), historical record quantile in the year of record occurrence (e-h), 2015 state likelihood (j-l), and PCC in 2050 (m-p). All patterns are from a single member randomly selected from the ACCESS-ESM1-5 ensemble. Shape parameter and historical record quantiles are computed based on GEV fits following the method indicated above the columns. GEV for future quantile evolution are in all cases based on the full ensemble, as in the main text, see 4 for details.



SI Figure S4: As main Fig. 3d-e, but with empirical (non-parametric) CDF per quantile level bin (on right y-axis), showing fraction of gridcells subject to the corresponding probability ratios

1426 S4 Supplementary figure to main Fig. 3d-e

1427 S5 Derivation quantile level of records

1428 Intuitively, the expected quantile level of the maximum M_j of an i.i.d. sample $\{X_1, \dots, X_j\}$ of length j
 1429 should correspond to $1 - \frac{1}{j+1}$. This follows from the fact that the marginal record breaking probability
 1430 at timestep j is $\frac{1}{j}$. For example, for timestep $j = 3$, the average record breaking rate is $\frac{1}{3}$, meaning that
 1431 the quantile level of the current record at time $j = 2$ is $1 - \frac{1}{3}$. This means that the record set at time
 1432 $j = 1$ has an average quantile level of 0.5, hence, $1 - \frac{1}{j+1}$.

1433 Formally, we derive this result as follows. We are looking for the quantile level of the maximum

¹⁴³⁴ M_j , i.e. $F_X(M_j)$, where F_X is the CDF of each of the i.i.d. data points X_i and M_j is the maximum
¹⁴³⁵ of X_1, \dots, X_j . In order to find the expected value of $F_X(M_j)$, we need an expression for the PDF of
¹⁴³⁶ $F_X(M_j)$, which we find by determining the CDF of $F_X(M_j)$ in equation (11) and taking the derivative
¹⁴³⁷ of that expression in equation (12).

$$F(F_X(M_j)) = P(F_X(M_j) \leq F_X(m)) = \\ P(M_j \leq m) = P(\max(X_1, \dots, X_j) \leq m) = P(X_1 \leq m) \cdot \dots \cdot P(X_j \leq m) = F_X(m)^j \quad (11)$$

¹⁴³⁸ In the equation above we use the property that $P(F_X(M_j) \leq F_X(m)) = P(M_j \leq m)$. To find the
¹⁴³⁹ PDF for $F_X(M_j)$, we take the derivative of the previous expression:

$$f(F_X(M_j)) = jF_X(m)^{j-1}f_X(m) \quad (12)$$

¹⁴⁴⁰ Now we can determine the expected value of $F_X(M_j)$ as follows:

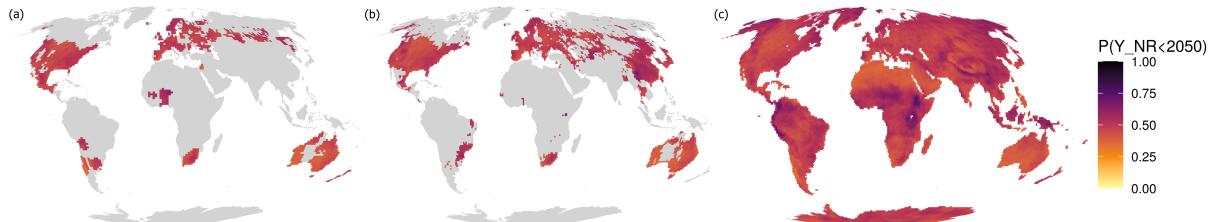
$$\mathbb{E}[F_X(M_j)] = \int F_X(m)f(F_X(M_j))dm = \int F_X(m)jF_X(m)^{j-1}f_X(m)dm = j \int F_X(x)^j f_X(x)dx \quad (13)$$

¹⁴⁴¹ We can substitute $u = F_X(x)$ in the previous integral and instead of integrating $f_X(x)dx$ we then
¹⁴⁴² integrate over du from 0 to 1, leading to the final result:

$$\mathbb{E}[F_X(M_j)] = j \int_0^1 u^j du = j \cdot \left[\frac{1}{j+1} u^{j+1} \right]_0^1 = \frac{j}{j+1} = 1 - \frac{1}{j+1} \quad (14)$$

¹⁴⁴³ Note: the above holds if $\{X_1, \dots, X_j\}$ is a stationary, i.i.d. sample. This property manifests in the
¹⁴⁴⁴ stationary CCP/MCP ratios in Fig. 3d-e.

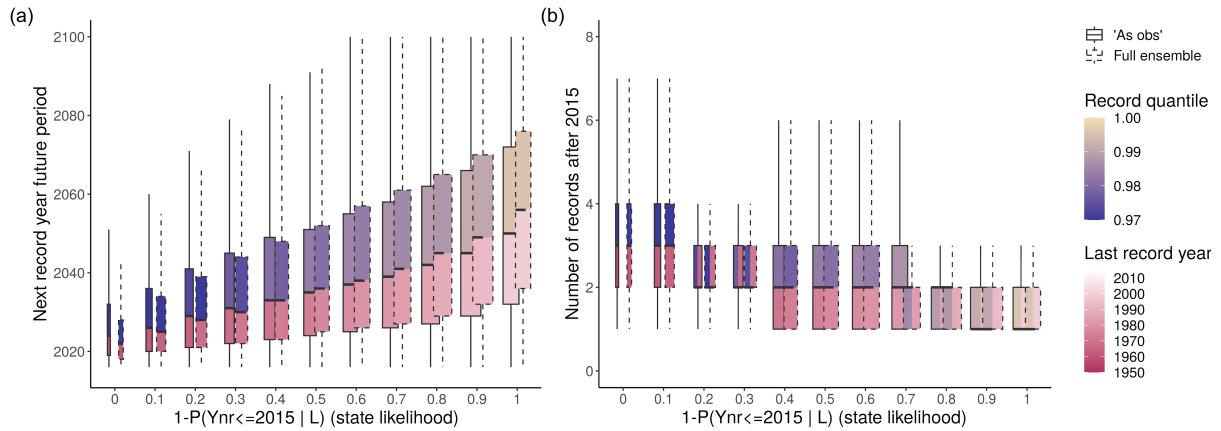
¹⁴⁴⁵ S6 Marginal cumulative record breaking probability patterns



SI Figure S5: Marginal cumulative record breaking probability (MCP) by 2050 as defined in the main text for HadEX3 (a), REGEN (b) and ERA5 (c).

¹⁴⁴⁶ Figure S5 shows the MCP by 2050 for the observational and reanalysis datasets. The MCP includes
¹⁴⁴⁷ only the effect of climate change on record breaking probabilities, and is independent of the historical
¹⁴⁴⁸ record level to be exceeded. The difference in magnitude between HadEX3 and REGEN on the one
¹⁴⁴⁹ hand and ERA5 on the other hand is due to the start year; for HadEX3 and REGEN the cumulative
¹⁴⁵⁰ summation starts in 2016, for ERA5 in 2024.

1451 **S7 Additional validation state likelihood**

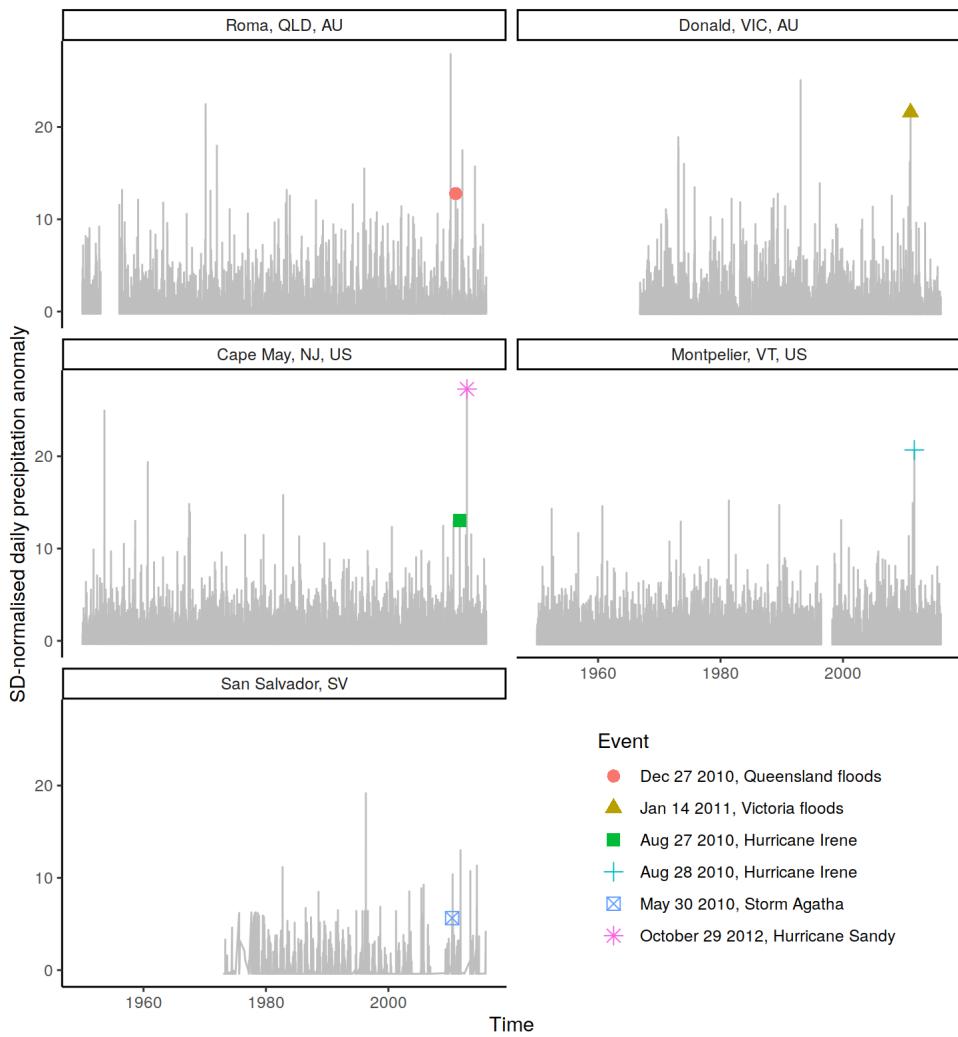


SI Figure S6: Correlation of the state likelihood (binned) in 2015 with the year of next record occurrence (a) and with the number of records in the period 2016–2100 (B), evaluated in all CMIP6 models. The colour shading of the bars show the associated bin-means of record quantile level and record-setting year T , with a clear gradient towards lower record quantile levels (less extreme) and longer-ago years as the state likelihood decreases.

1452 Fig. S6 shows the correlation between state likelihood and indicators of future record breaking in
 1453 CMIP6 models, in part validating its use as an indicator of disaster potential.

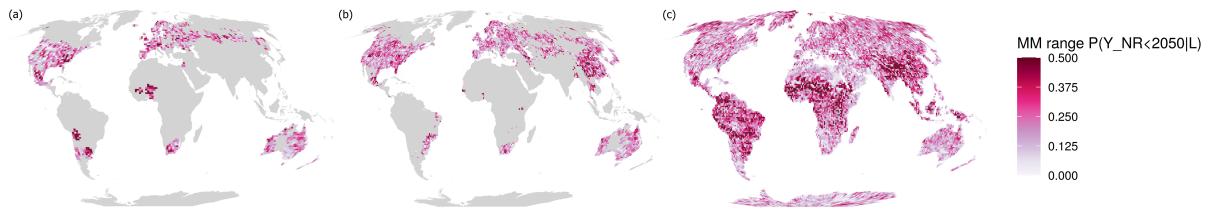
1454 Fig. S7 shows that the selected single station data in the regions discussed in the main text shows
 1455 a clear signature corresponding to the events we associated with the record breaking. These stations
 1456 were selected based on their location being in the gridcells of interest; gridcells with 2009 state likelihood
 1457 ≤ 0.25 and record breaking in the years 2010–2015, and reported in the disaster reports of the events
 1458 in question [57–60]. Not all stations show the maximum daily event, which is expected as HadEX3
 1459 aggregates multiple stations in their Rx1d product, but all stations show daily precipitation values in
 1460 the uppermost quantiles of the full sample. Storm Agatha (San Salvador) is least well represented, also
 1461 in other stations in the region that we assessed. The lower data quality in Central America plays a big
 1462 part (exemplified by the sparsity of the record for San Salvador): this impairs confidence in both the
 1463 accuracy of records in HadEX3 as well in the verification data itself.

1464 For reference, we added a marker for Hurricane Sandy in the timeseries for Cape May, which evidently
 1465 led to record-breaking precipitation in New Jersey, where it made landfall. This event of October 2012
 1466 is not clearly visible as record breaking cluster in HadEX3, which we hypothesise could be associated
 1467 with the single landfall of Sandy, as opposed to the repeated ‘bouncing’ landfall of Irene.



SI Figure S7: Daily precipitation observations from single stations from the GHCN-Daily network [90–93] and the Australian Bureau of Meteorology [94, 95]. Events mentioned in the main text are indicated with markers [57–60].

1468 **S8 Model uncertainty in probability estimates**



SI Figure S8: Intermodel range (difference between maximum and minimum projected value) for CCP 2050. The spread is due to model uncertainty in the temporal evolution of the record quantile, thus a combined measure of differences in climate sensitivity and local patterns of Rx1d changes

1469 Fig. S8 shows the model uncertainty in record breaking projections – the range of the model-mean
1470 ensemble of CCP 2050 projections is shown. The overarching pattern is the well-known uncertainty
1471 pattern where precipitation and/or Rx1d changes in the tropics and monsoon regions are most uncertain
1472 and feature highest intermodel differences. These are also the regions that are projected to see the
1473 largest changes in both absolute and relative magnitude of extreme precipitation [52], and are also most
1474 vulnerable. The pattern is modified by the local record quantiles; where CCP values are largest due to
1475 the combination of low record quantiles and strong climate change, the uncertainties are largest too.