

SUPPLEMENTARY INFORMATION

Off-fault damage controls near-surface rupture behaviour in soft sediments

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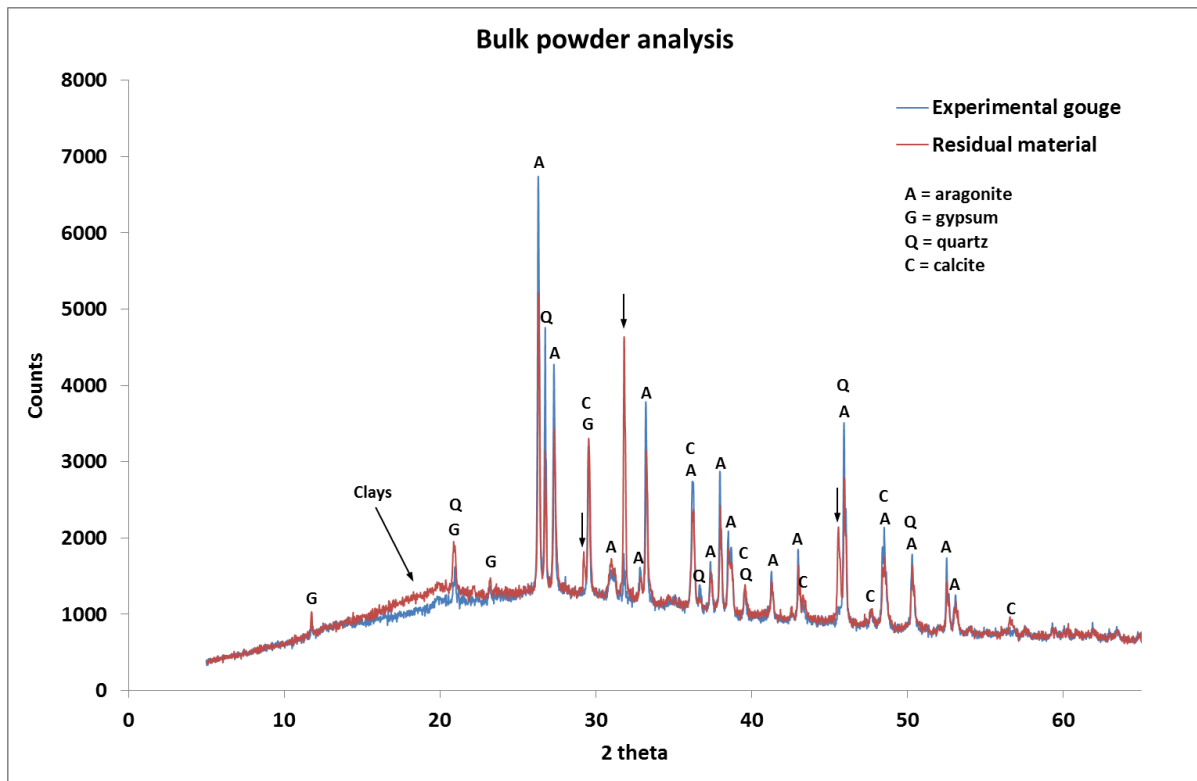
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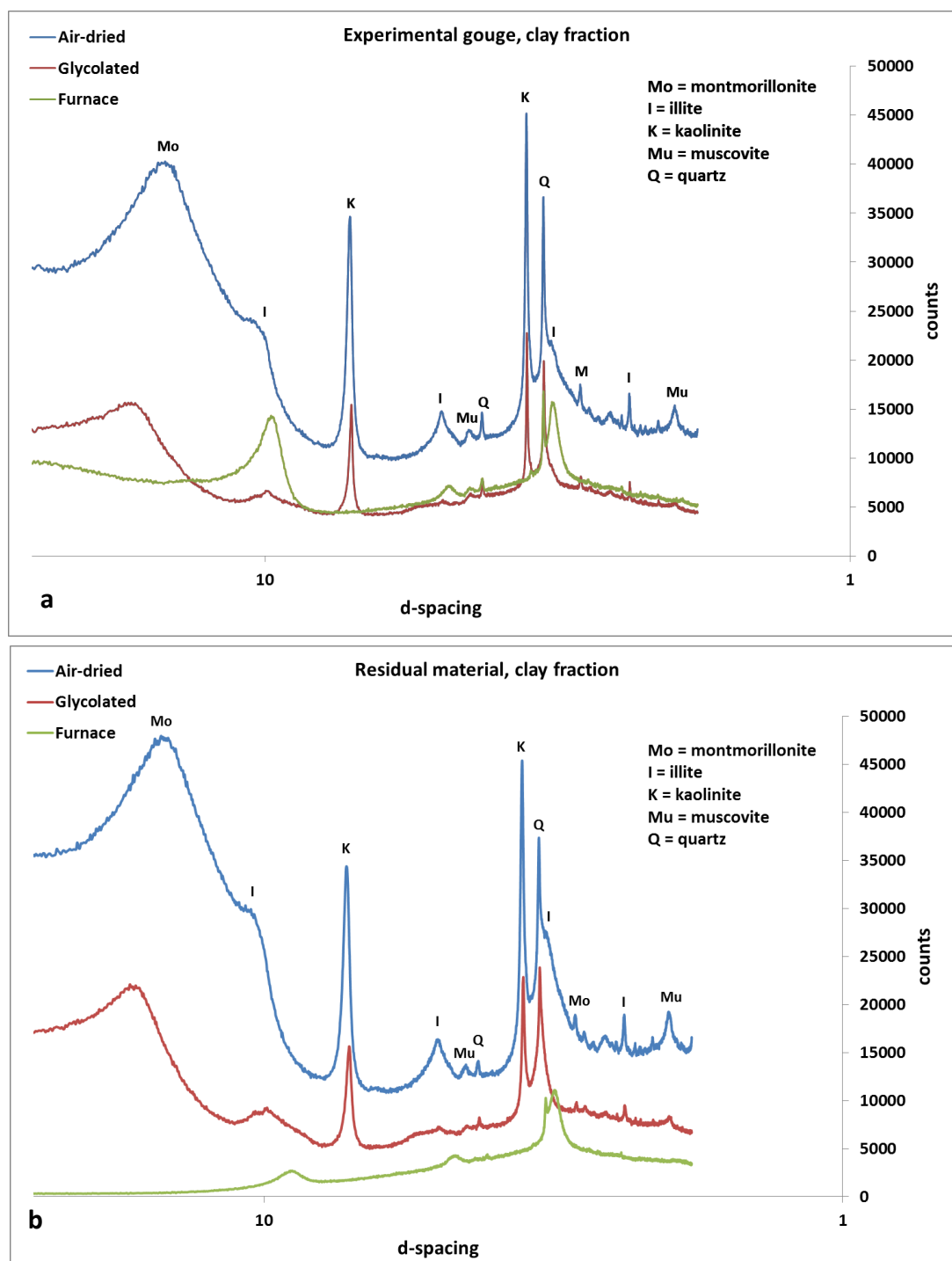
Supplementary Information I – Lisan Formation and experimental simulated gouge composition

Synthetic gouges were prepared for testing by disaggregating undeformed samples of the Lisan Formation sediment collected in the field. We passed the disaggregated material through a 1 mm sieve, to remove anomalously large grains or aggregates, such as late-stage mineralisation crystals (predominantly gypsum) and fragments of fossil wood, and to ensure that all experiments had a reasonably consistent starting grain size. X-Ray diffraction (XRD) analyses of the material which passed through the sieve, and of the residual material that did not, show that their compositions are almost identical and thus the sieving process did not result in a biased sampling of finer-grained components of the sediment (Supplementary Figure 1).

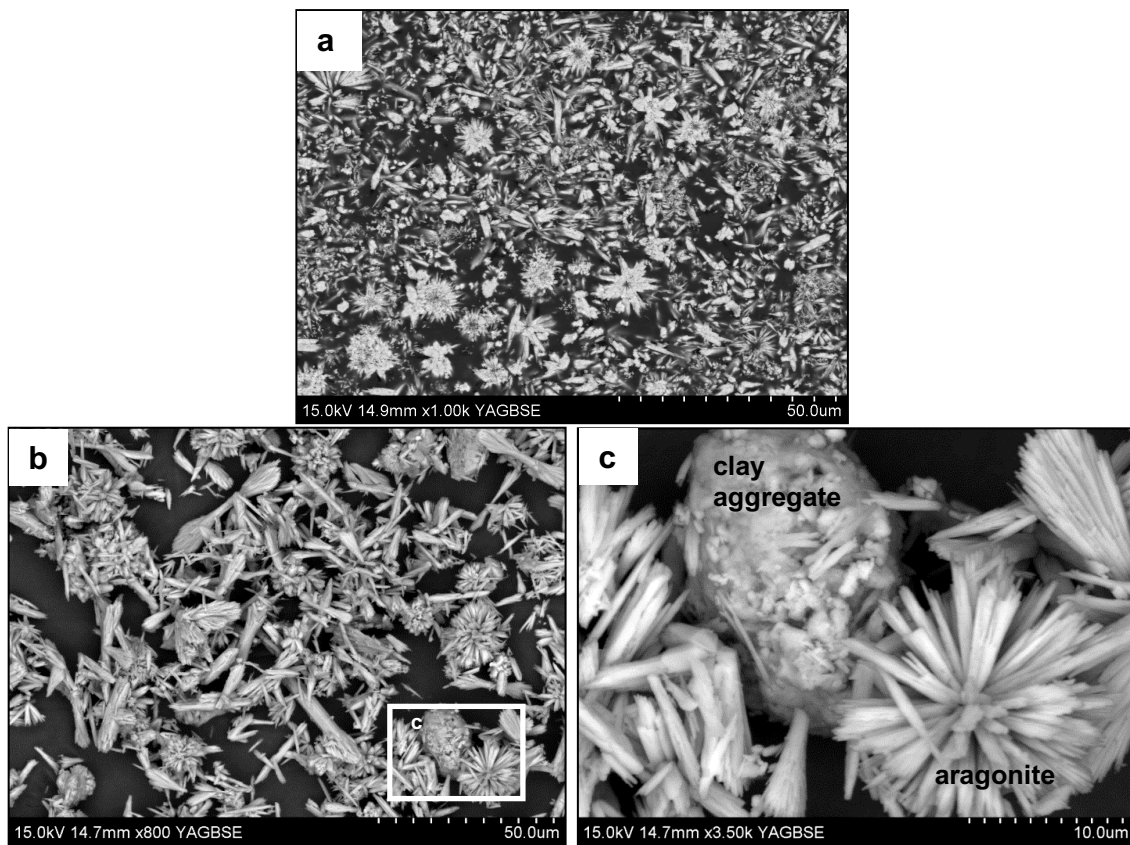
The gouge is composed predominantly of aragonite and clays, which comprise between 40-60% and 30-50% of the deformed gouges, respectively, plus small amounts (up to 10%) of gypsum, quartz and calcite (Supplementary Figure 1). Separate XRD analysis of the clay fraction (<2 μm grain size) shows that it comprises mixed layer illite-smectite and kaolinite (Supplementary Figure 2). The aragonite crystals, which comprise the majority of clasts within the gouge, have a fine grain size of <15 μm and distinctive acicular habits (Supplementary Figure 3a), which are sometimes radially arranged in rosette-type structures (Supplementary Figure 3b-c). The less frequent clasts of gypsum, quartz and calcite have grain sizes up to 100 μm . The reason for using a relatively coarse 1 mm sieve was to minimise disruption and damage to individual grains and aggregates. Scanning electron microscope images of the sieved gouge confirm that the acicular aragonite crystals are un-damaged by sieving and retain their radial arrangements (Supplementary Figure 3b-c); they therefore still closely resemble the natural state of the Lisan Formation material (Supplementary Figure 3a).



Supplementary Figure 1: XRD patterns of the sieved gouge (<1 mm) used in the experiments and the residual material, which did not pass through the sieve. The compositions are almost identical, with the exception of the arrowed peaks, which are relatively enriched in the residual material. Possible matches of these peaks are barite and basanite. Either of these would be consistent with the residual material being enriched in late-stage crystalline mineralisation products, which one would expect to be coarser and more resistant to disaggregation than the rest of the Lisan sediment. The broad ‘bump’ present at low values of 2 theta is typical of clay minerals, which are poorly crystalline and so do not produce strong, clear diffractions. To identify the clays, we separated the clay fraction to produce oriented mounts (see Supplementary Figure 2).

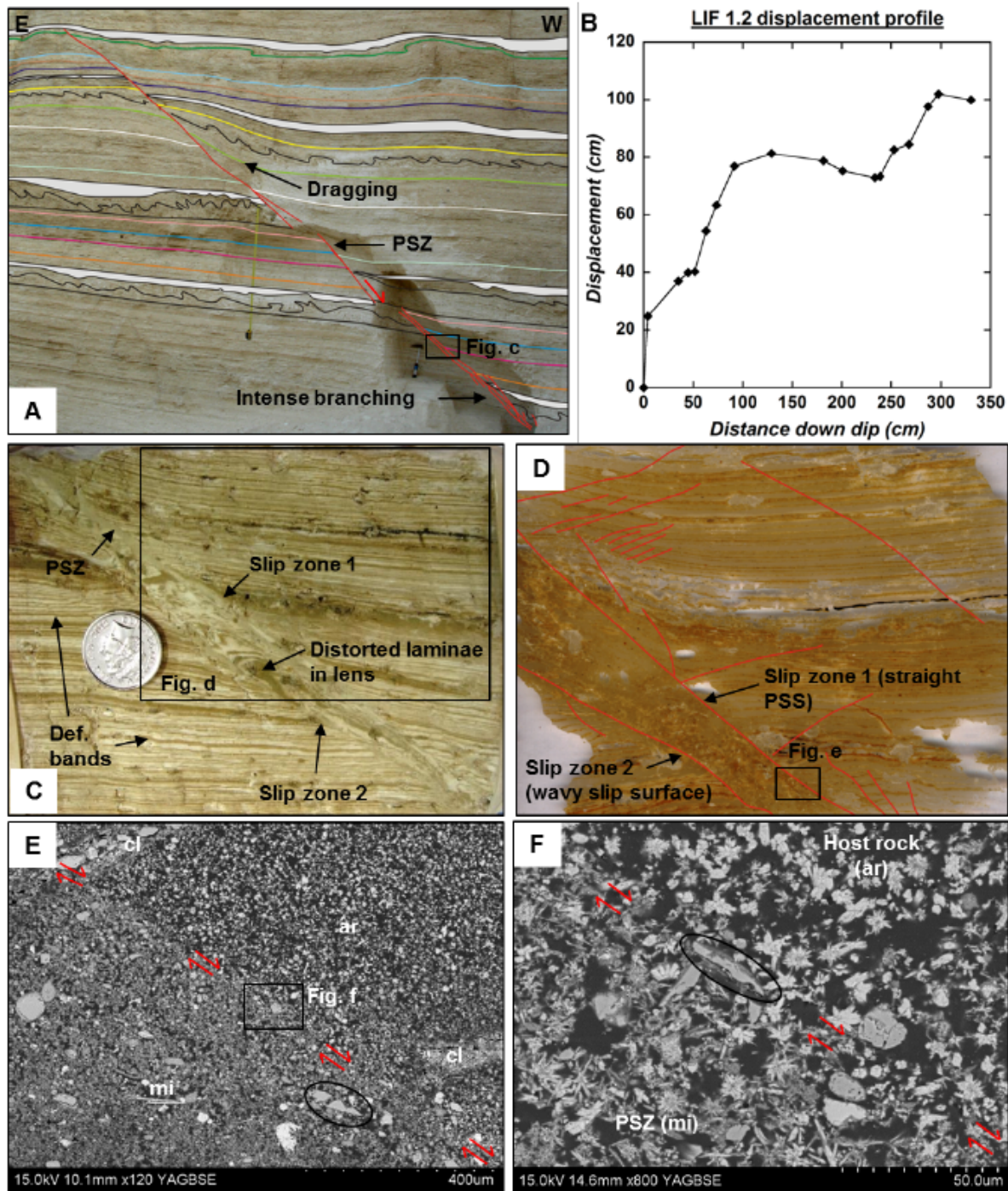


Supplementary Figure 2. **a)** XRD patterns of the air-dried, glycolated and furnace-treated (550°C) clay fraction of the sieved Lisan sediment used in the experiments. **b)** XRD patterns of the air-dried, glycolated and furnace-treated (550°C) clay fraction of the residual material, which did not pass through the sieve, and so was not used in the experiments. The composition is the same as that observed for the experimental gouge, clay fraction shown in panel **a**.



Supplementary Figure 3 – Microstructures of the natural protolith and the sieved simulated gouges used in the friction experiments. *a)* SEM image of an aragonite layer from an undeformed and undisturbed sample of the Lisan Formation. The aragonite crystals are acicular with sharp, pointed terminations, and sometimes arranged in stellate or radiating aggregates, resembling rosettes. *b)* SEM image of the sieved gouge used for the experiments. *c)* Higher magnification image of the boxed area in panel **b**. It is evident from both panels **b** and **c** that radial arrangements of acicular aragonite crystals have retained the structures observed in the natural sample shown on panel **a**, so they have not been disturbed or damaged during the disaggregation and sieving process.

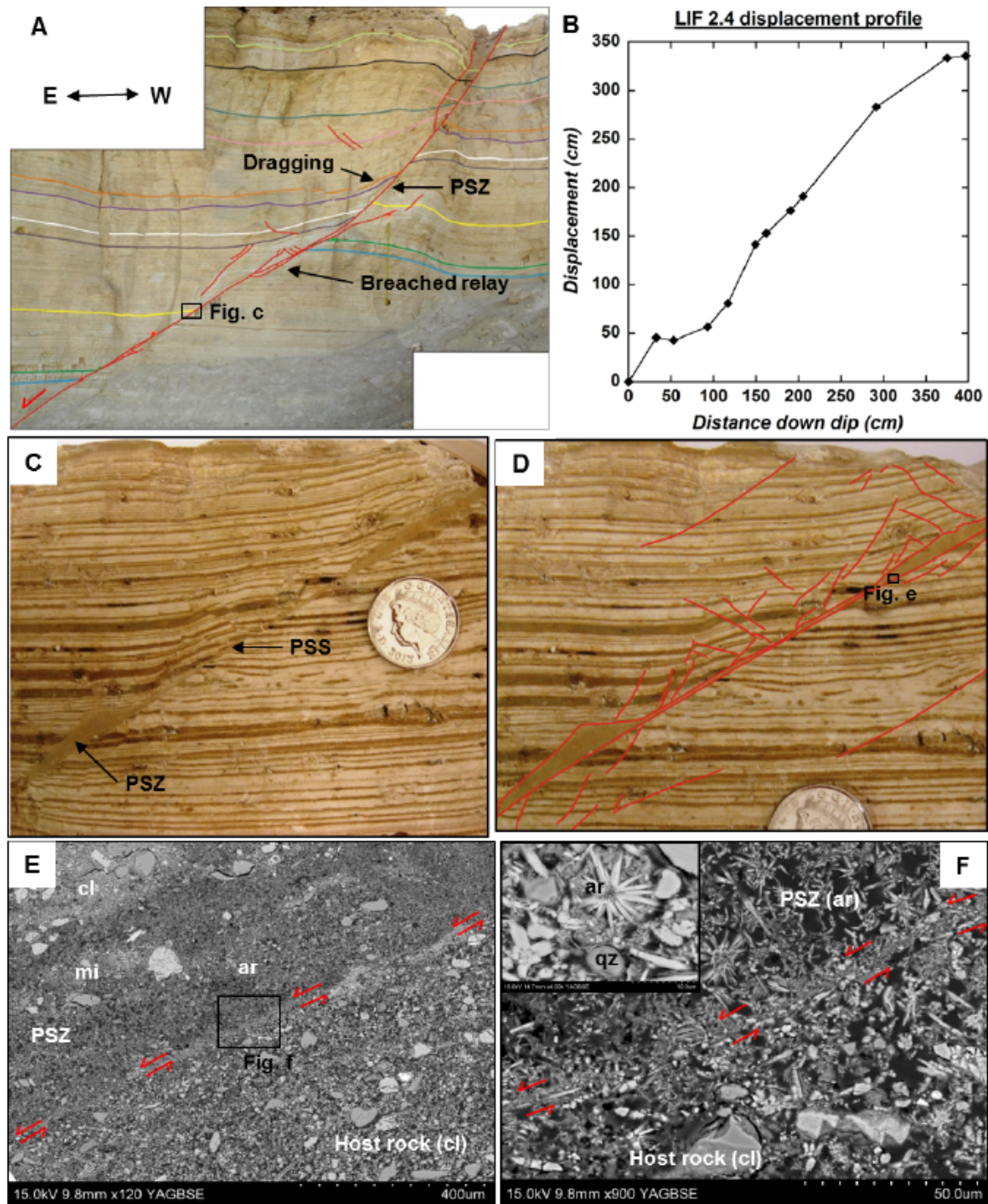
Supplementary Figure 4 – The Masada fault zone (MFZ) structure



Supplementary Figure 4 – Structure and microstructures of the syn-sedimentary Masada Fault Zones. **A)** Fault LIF. 1.2, $d_{max} = 102$ cm, viewed at the meso-scale in outcrop. Mixed layers are shaded grey and slumps marked in black. **B)** Displacement profile constructed for fault LIF 1.2. The plotted points correspond to the highlighted sedimentary layers in Fig. A. **C)** Hand sample from fault LIF 1.2. The location from which it is taken is shown in Fig. A. The displacement at this point is about 98 cm. **D)** High-resolution scan of thin section

taken from boxed area in Fig. C. The slip surface bounding slip zone 1 is observed to be straighter than that bounding slip zone 2, thus we infer it to be the principal slip surface SSS. **E)** SEM image of the PSS. The location from where the image is taken is shown in Fig. D. The location of the PSS is indicated by shear-sense arrows; its location is discernible due to the subtle difference in composition between the host rock, which is predominantly aragonite, and the PSZ material, which is more clay rich. A group of aligned clasts close to the PSS is circled, ar = aragonite, cl = clay-rich clastic material, mi = areas where the aragonite and clay-rich materials have mixed. **F)** Close up image of PSS and PSZ shown in Fig. E, showing the similarity in grain size of the material within the PSZ compared to the host rock. Several aragonite crystals within the slip zone maintain their concentric structures. Again some alignment of clasts along the PSS is evident (circled).

Supplementary Figure 5 – The Masada fault zone (MFZ) structure



Supplementary Figure 5 – Structure and microstructures of the syn-sedimentary Masada Fault Zones. **A)** Fault LIF 2.4, $D_{max} = 337$ cm, viewed at the mesoscale in outcrop. **B)** Displacement profile constructed for fault LIF 2.4. The plotted points correspond to the highlighted sedimentary layers in Fig. A. **C)** Hand sample from fault LIF 2.4. The location from which it is taken is shown in Fig. A. The displacement at this point is about 280 cm. **D)** An annotated zoomed in portion of Fig. C is shown to highlight PSZ boundaries and small scale

off-fault damage. Again, one PSZ boundary (labelled PSS) is observed to be much straighter than the opposite PSZ boundary. **E)** SEM image of the PSS. The location from where the image was taken is shown in Fig. D. The location of the PSS is indicated by shear-sense arrow; its location is discernible due to the subtle difference in composition between the host rock and the PSZ material, and an enrichment of clay along the PSS due to clay smear. The PSZ is very heterogenous in composition, e.g. note the clump of clay-rich clastic material labelled in the top left-hand side of the image, compared to the rest of the PSZ, which is aragonite rich. ar = aragonite, cl = clay-rich clastic material, mi = areas where the aragonite and clay-rich materials have mixed. **F)** Close up image of PSS and PSZ shown in Fig. E; location of PSS is indicated by shear-sense arrows and highlighted by clay smear and an alignment of acicular aragonite needles. Note how aragonite crystals in the PSZ retain their concentric arrangements of needles, showing little evidence for fracturing or comminution. Inset shows an example of an aragonite crystal and quartz (qz) clasts within the PSZ, which are surrounded by a cortex of clay.

Supplementary Information II: Experimental apparatus specifications and sample assembly geometry.

High velocity rotary shear apparatus

Friction experiments were performed using the low to high velocity rotary shear apparatus (model MIS-233-1-77, built by the Marui & Co., Ltd. Company, Osaka, Japan) at the Rock Mechanics Laboratory, Durham University (Supplementary Figure 6). The apparatus is housed in a rigid loading frame, made of assembled steel plates, and set up in a vertical configuration, secured to the floor by a base plate (Supplementary Figure 6a-b).

Samples are mounted onto sample holders (Supplementary Figure 6c), where they are kept fixed to the main shafts and axially aligned during the experiments by a hydraulic mechanical lock (Supplementary Figure 6c). During the experiments, the upper main shaft rotates, whilst the lower one remains stationary. The upper main shaft of the apparatus is driven by an electric, servo-controlled motor with 11 kW power, 70 Nm rated torque and a maximum rotation speed of 1500 RPM. The revolving speed of the main shaft ranges from 1.5×10^{-6} -1500 RPM, and is controlled using a system of clutches and reduction gears located in the gearbox (Supplementary Figure 6a-b). This corresponds to a peripheral slip rate of 6 cm/yr to 2 m/s, when applied to cylindrical samples with an outer diameter of 25 mm. The apparatus can accelerate to the maximum peripheral target speed of 2 m/s in 205-270 ms at normal stresses of 3 and 18 MPa, respectively, with accelerations of 3.7 and 4.9 m/s², respectively.

An axial load is applied to the lower main shaft by a pneumatic piston (Bellofram type cylinder) with an 82 mm stroke, and it is able to produce a thrust of 10 kN (Supplementary Figure 6a-b). The axial load system is equipped with a high-precision air regulator to reduce and automatically correct the load fluctuation during loading. Axial load measurements are performed using a thin compression load cell (strain gauge type, Supplementary Figure 6a-b) with a 10 kN capacity and a rated output of 2 mV/V $\pm 0.5\%$. Axial load cell resolution is ± 0.005 kN.

Torque values attained during the experiments are measured by two compression load cells (strain gauge type) (Supplementary Figure 6a-b), which are activated by a torque bar fixed to the lower main shaft. The load cell capacity is 1 kN, with a rated output of 2 mV/V $\pm 0.5\%$. Torque cell resolution is $\pm 5 \times 10^{-4}$ kN.

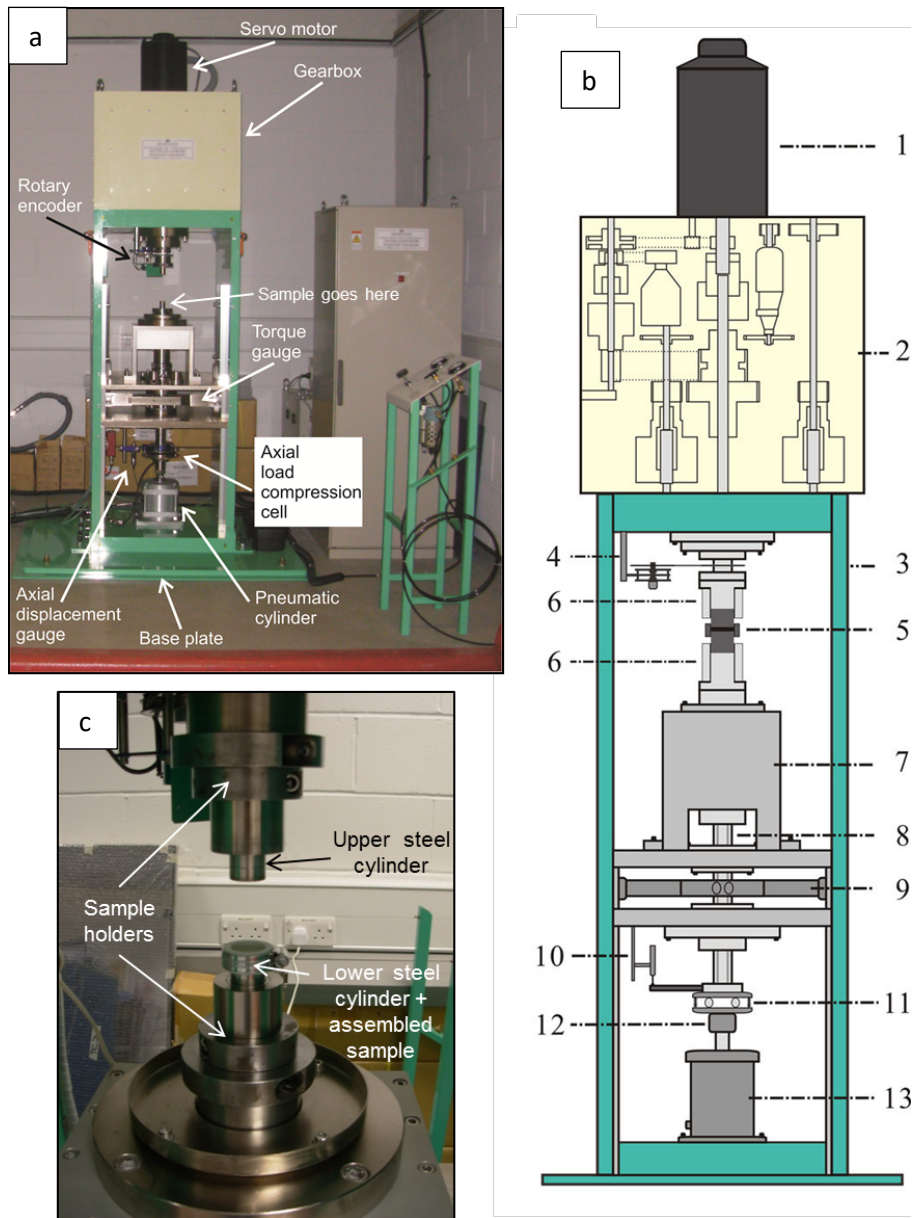
The axial displacement values attained during the experiments are measured using a high sensitivity displacement gauge (strain gauge type) with a 10 mm capacity and a rated output of 5 mV/V $\pm 0.1\%$ (Supplementary Figure 6a-b). Axial displacement resolution is $\pm 2 \times 10^{-3}$ mm.

The revolution speed and the cumulative number of revolutions are measured by a tachometer and a pulse counter, respectively, based on the pulses measured by a rotary encoder with a capacity of 3600 pulses/revolution (Supplementary Figure 6a-b).

Sample assembly

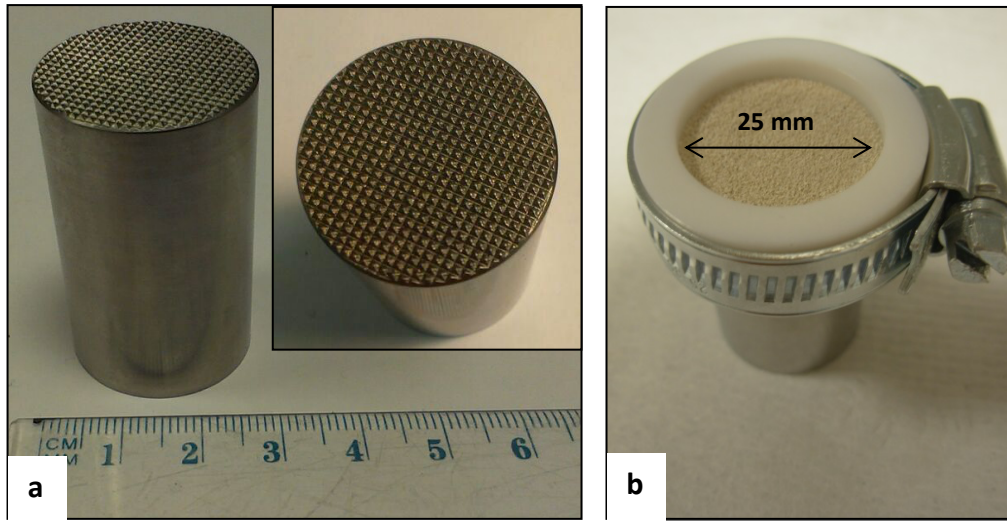
A synthetic fault zone was created by sandwiching up to 1.4 g of disaggregated Lisan Formation sediment between two stainless steel cylinders (25 mm in diameter), whose ends are machined to a depth of 500 μm in a gridded fashion (Supplementary Figure 7), to simulate slip surface roughness and to ensure that slip localises within the gouge layer, rather than at the boundaries between the gouge and the steel cylinders. To limit gouge loss during the experiments, the sample assembly was confined using a Teflon ring¹, which was tightened to the steel cylinder using a hose clip (Supplementary Figure 7b). The inner edges of the Teflon rings are machined to reduce their sharpness and avoid ring damage and sample contamination by Teflon.

¹ Mizoguchi, K., Hirose, T., Shimamoto, T., Fukuyama, E. Reconstruction of seismic faulting by high-velocity friction experiments: An example of the 1995 Kobe earthquake. *Geophys. Res. Lett.* **34**, L01308 (2007).



Supplementary Figure 6. **a)** Photograph of the low to high velocity rotary shear apparatus installed in the Rock Mechanics Laboratory at Durham University. **b)** Close up image showing an assembled sample mounted in the apparatus. **c)** Schematic diagram showing the main units of the apparatus, from Ma et al. (2014)². 1 = servo motor, 2 = gear box, 3 = loading frame, 4 = rotary encoder, 5 = specimen assembly, 6 = locking devices holding specimen assembly, 7 = frame for holding the lower loading column, 8 = axial loading column, 9 = torque gauge, 10 = axial displacement transducer, 12 = axial force gauge, 13 = air actuator.

² Ma, S. et al. A rotary-shear low to high-velocity friction apparatus in Beijing to study rock friction at plate to seismic slip rates. *Earthq. Sci.* **27**, 469-497 (2014).



Supplementary Figure 7. a) Stainless steel cylinders used during the experiments, whose ends are machined in a gridded fashion. **b)** Assembled gouge sample.

Experimental determination of slip rate and slip for cylindrical samples

To calculate the slip rate within the gouge layer, it is necessary to convert angular velocity, ω (in rad/s; RPM can be converted to rad/s by multiplying by $2\pi/60$) to linear velocity, v (in m/s). This can be done using the equation:

$$v = \omega r \quad (\text{Eq. 1}),$$

where r is the sample radius (in metres). Also, given the cylindrical shape of the samples, v increases with sample radius. Thus, the slip rate that we use is an “equivalent slip rate”^{3,4}, V_e , corresponding to the slip rate within the sample at $2/3r$. Taking into account the above factors, the equation for calculating V_e is:

$$V_e = \frac{4\pi Rr}{3 \times 60} \quad (\text{Eq. 2}),$$

where R is the revolution rate of the motor in RPM. We refer to the equivalent slip rate simply as the slip rate v in this thesis and, as a consequence, the, displacement, d is calculated as

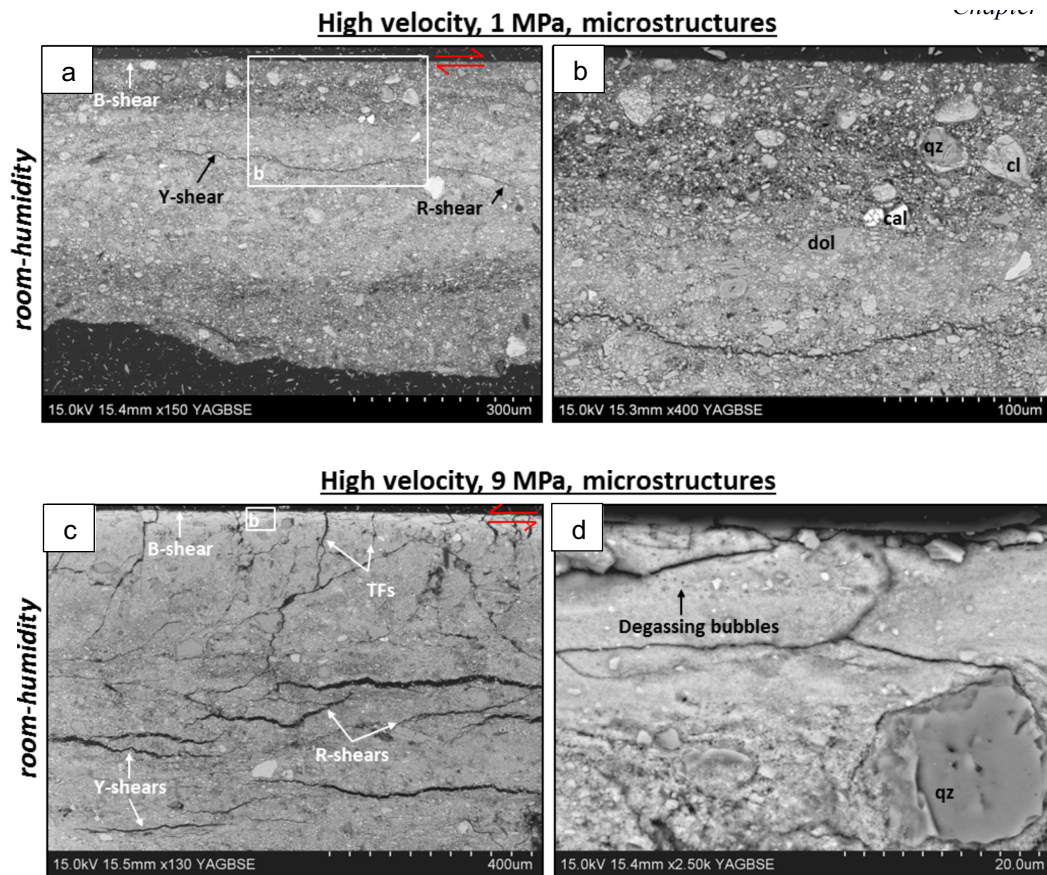
$$d = V_e t \quad (\text{Eq. 3}),$$

where t is the time.

³ Hirose, T. & Shimamoto, T. Growth of molten zone as a mechanism of slip weakening of simulated faults in gabbro during frictional melting. *J. Geophys. Res-Sol. Ea.* **B05202**, doi:10.1029/2004JB003207 (2005).

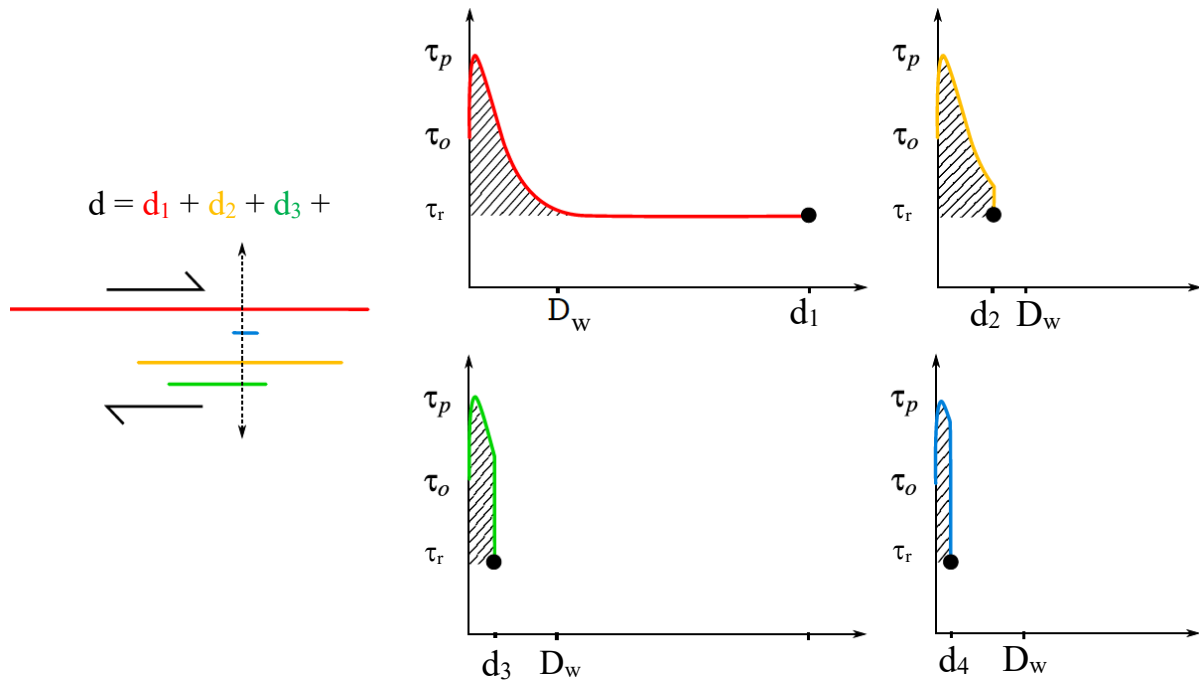
⁴ Shimamoto, T. & Tsutsumi, A. A new rotary-shear high-speed frictional testing machine: Its basic design and scope of research (in Japanese with English abstract). *Journal of the Tectonic Research group of Japan* **39**, 65-78 (1994).

Supplementary Figure 8 – Microstructures of seismic, high velocity experiments



Supplementary Figure 8. Microstructures of gouges deformed at seismic slip velocity (1.3 m s^{-1}) under 1 - 9 MPa normal load and at room-humidity conditions. Larger clasts are labelled to aid mineral identification: cal = calcite, cl = clayrich aggregate, dol = dolomite, qz = quartz. Areas labelled 'R' are resin. **a)** Room-humidity gouge (experiment 532, $d = 36.77 \text{ m}$). **b)** High magnification image of the boxed area in fig. a. **c)** Room-humidity gouge (experiment 624, $d = 4.6 \text{ m}$). TFs = tensile fractures. **d)** High magnification image of the boxed area in fig. c.

Supplementary Figure 9 – Energy dissipation during different amounts of fault slip



Supplementary Figure 9. The area of the triangle, highlighted under the slip-weakening curve shown in red, represents the energy dissipated in the creation of a single shear fracture with an amount of slip $d_1 \geq D_w$. The areas of the trapezoid, highlighted under the slip-weakening curves shown in yellow (d_2), green (d_3) and blue (d_4), represent the energy dissipated in the creation of a single shear fracture with an amount of slip $d < D_w$.