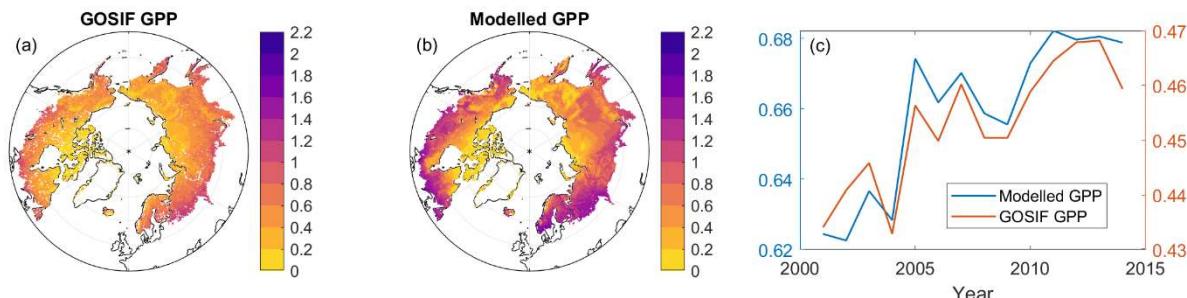


1 Extended Data

2 Model evaluation for different historical periods

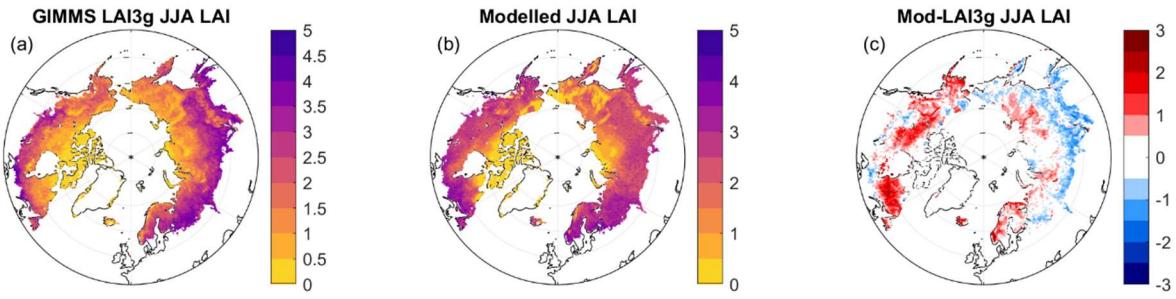
3 For the historical period (1901-2014) using observation-based climate data, the modelled leaf area index (LAI),
4 gross primary productivity (GPP) and aboveground carbon (ABC) are evaluated against satellite-based dataset
5 (Extended Data Figs.1-3). The modelled ecosystem-level BVOC emissions are evaluated with observations from
6 the available literature (Extended Data Table 1).

7 The modelled gross primary productivity (GPP) has been evaluated against the OCO-2-based SIF GPP product
8 (GOSIF)¹, and the mapping of multiple-year averages show general agreements in spatial patterns (Extended
9 Data Fig. 1a, b). The modelled GPP is higher than estimates based on GOSIF in the southern boreal forest and
10 also in terms of areal average. The modelled interannual variability of annual areal averaged GPPs correlates
11 well with the observations based on GOSIF (Extended Data Fig. 1c).



12
13 *Extended Data Figure 1. Comparing the modelled annual gross primary productivity (GPP) with the GOSIF GPP product.*
14 *(a) Annual GOSIF-based GPP averaged over the period of 2001-2014; (b) LPJ-GUESS modelled annual GPP averaged over*
15 *the period of 2001-2014; (c) the timeseries of annual, areal average of GPP (KgC/m²) over the study domain (including*
16 *tundra and boreal regions).*

17
18 The modelled LAI averaged over June, July and August are compared with estimates from the GIMMS LAI3g²
19 product over the same period (Extended Data Fig. 2). We find that the modelled spatial patterns of LAI are
20 similar to GIMMS LAI3g though the absolute differences (Extended Data Fig. 2c) show that the model tends to
21 overestimate LAI in some of tundra regions, such as in north America, Norway and northern parts of Finland and
22 small regions of Siberia. In contrast, slight underestimations are found in the southern part of boreal region.

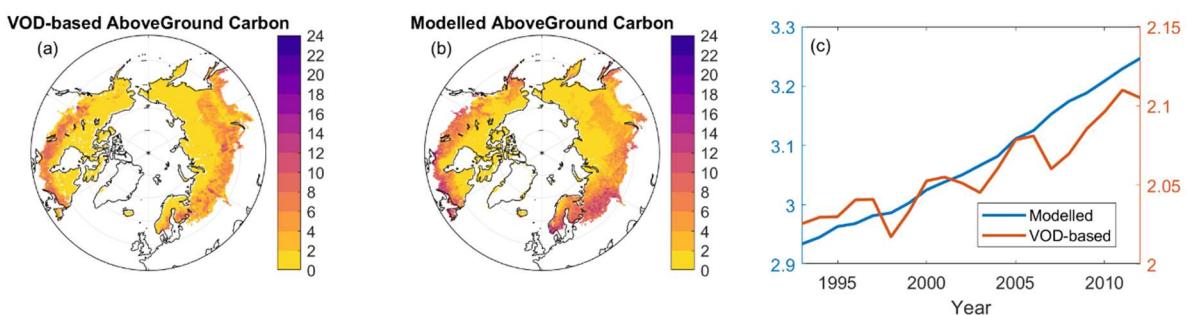


23

24 *Extended data Figure 2. Comparing the modelled leaf area index (LAI) over the growing season (June, July and August, JJA)*
 25 *with the GIMMS LAI3g product over the period 1982-2011. (a) GIMMS LAI3g-observed averaged LAI in JJA; (b) LPJ-*
 26 *GUESS modelled LAI in JJA; (c) the difference in the modelled JJA LAI between the modelled and the GIMMS LAI3g*
 27 *product.*

28

29 The modelled aboveground carbon pool (i.e., leaf and stem carbon in vegetation) is compared with estimates
 30 from a vegetation optical depth (VOD)-based product³. The overall spatial patterns are captured by the model
 31 (Extended Data Fig. 3a, b), and though there is an overestimation of areal averages of aboveground carbon from
 32 our model, the increasing trends are well represented (Extended Data Fig. 3c).



33

34 *Extended Data Figure 3. Comparing the modelled annual aboveground biomass carbon (ABC, KgC m⁻²) with vegetation*
 35 *optical depth (VOD)-based ABC product. (a) VOD-based ABC averaged over the period of 1993-2012; (b) LPJ-GUESS*
 36 *modelled aboveground carbon (including biomass from leaf and stem) averaged over the period 1993-2012; (c) the*
 37 *timeseries of areal averaged ABC over the study domain (including tundra and boreal regions).*

38

39 The above comparison of modelled and regional estimates of GPP, LAI and ABC show that the model can
 40 generally capture the spatial and temporal changes of these variables, though with some overestimations of GPP
 41 and ABC for southern boreal regions and of LAI for the tundra region.

42 The modelled ecosystem-level isoprene and monoterpene fluxes are compared with the published values from
 43 different ecosystems (Extended Data Table 1), which show that the model produce fluxes of similar magnitudes
 44 to the observed emissions. We note, however, that an exact comparison was not possible since the LPJ-GUESS
 45 model runs at a spatial resolution of 0.5 degrees, and although we used the output from the nearest gridcells with

46 a similar dominant vegetation type as observed at the sites, there could still have differences in terms of overall
 47 vegetation composition and microclimatic conditions.

48 *Extended Data Table 1 Ecosystem-level BVOC evaluation. The modelled values from the nearby gridcell were selected and*
 49 *the modelled units were converted to the same one as the one in the literature. Noted: only ecosystem-level observations were*
 50 *extracted from the literature. Dom.: dominated; MT: monoterpenes; ISO: isoprene*

Ecosystem types	Location	Dom. Species	Time	Compound s	Units	Observe d	Modelle d	Refs
Boreal forest	Siberian larch tree	<i>Larix cajanderi</i>	Jun, 2009 Jul, 2009	MT MT	mgC/m ² /d	3.3±2.9 2.4±1.6	2.0 2.4	⁴
	Scots pine	<i>Pinus sylvestris</i>	Mar, 2010-2013	MT	mgC/m ² /m	10.87	6.82	⁵
			Apr, 2010-2013			27.44	23.40	
			May, 2010-2013			85.08	54.99	
			Jun, 2010-2013			114.35	75.96	
			Jul, 2010-2013			163.07	111.15	
			Aug, 2010-2013			103.98	77.53	
			Sep, 2010-2013			57.18	42.97	
			Oct, 2010-2013			30.72	24.18	
			Nov, 2010-2013			6.63	15.73	
Tundra upland	Greenland tundra	<i>Cassiope tetragona</i>	Aug, 2009	ISO MT	µgC/m ² /h	1.38 21.54	0.014 0.014	⁶
	Alaska tundra	<i>Salix pulchra</i>	Jun-Jul 2005,2010&2011	ISO	µgC/m ² /h	Up to 1200	63.2	⁷
	Alaska tundra	Tussock tundra	Summer, 2018&2019	ISO MT	µgC/m ² /h	0.2-225 <1	0.2-118.8 0.34-4.81	⁸
	Dry hummocks	Shrub and mosses	July, 2007	ISO	µgC/m ² /h	24.5	85	⁹
	Boreal fen	<i>Sphagnum</i> mosses and sedges	July, 2007	ISO	µgC/m ² /h	186-220	198	¹⁰
	SubArctic fen	Graminoid	2006	ISO	µgC/m ² /h	Up to 1385	439	¹¹
	Tundra wetland	SubArctic fen	Jul, 2018 Jul, 2018	ISO MT	µgC/m ² /h	310 14	330 14	¹²

51

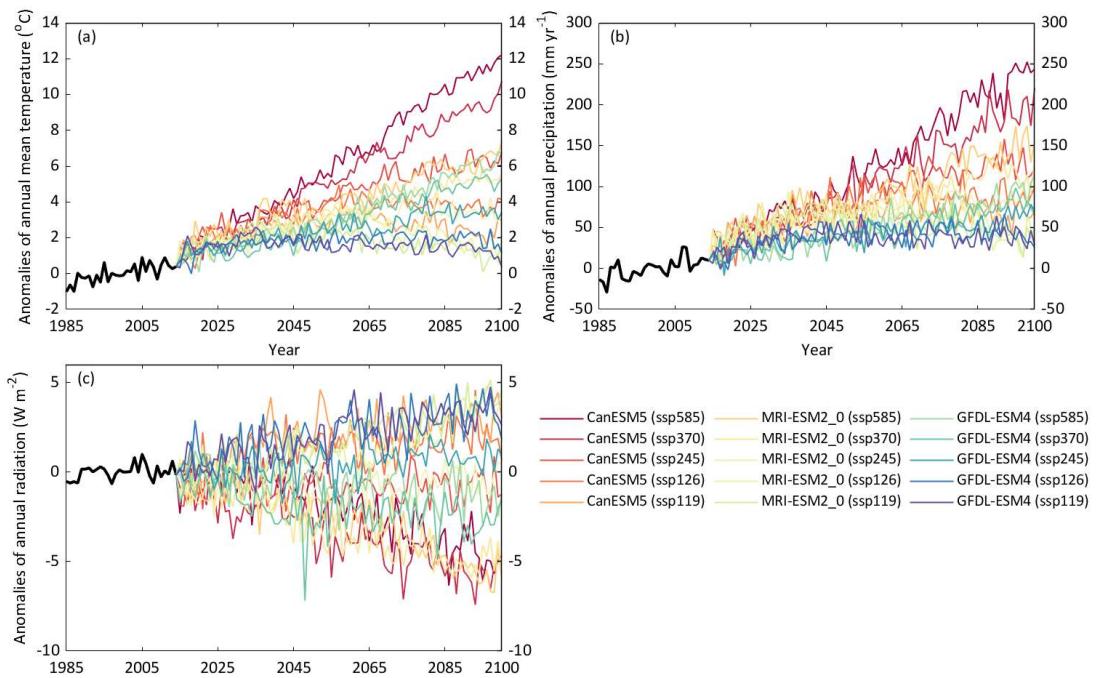
52 CMIP6 and predicted future climate change

53 We present details of the selected future emission scenarios and included general circulation model (GCM) in
 54 Extended Data Table 2. The design and description of the CMIP6 experimental protocol can be found in Eyring,
 55 et al.¹³ and the outputs from these models were downloaded through ESGF (<https://esgf-node.llnl.gov/search/cmip6/>) for the period 1901-2100 (Date of access: Sep-2020). The climatology of the CRU-
 56 NCEP and CMIP6 datasets over the period 1985-2014 was calculated, and the monthly biases between CRU-
 57 NCEP and each CMIP6 model data were calculated. The biases were corrected to each climate field for the
 58 whole future period (2015-2100). The bias-corrected temperature, precipitation and radiation were used to drive
 59 the LPJ-GUESS simulation over the future period. The anomalies in Extended Data Fig. 3 show that over the
 60 study region, the predicted temperature increase can be up to 12 °C, the annual precipitation increase can be up
 61 to 250 mm yr⁻¹ and the annual radiation show both increases and decreases. There is a general negative
 62 correlation between the changes in temperature and annual radiation. All of these 15 scenarios show that this
 63

64 region could become warmer and wetter, with a large range of responses between different CO₂ emission
65 scenarios.

66 *Extended Data Table 2 Overview of the selected future emission scenarios from CMIP6 and the general circulation model*
67 *included in this study.*

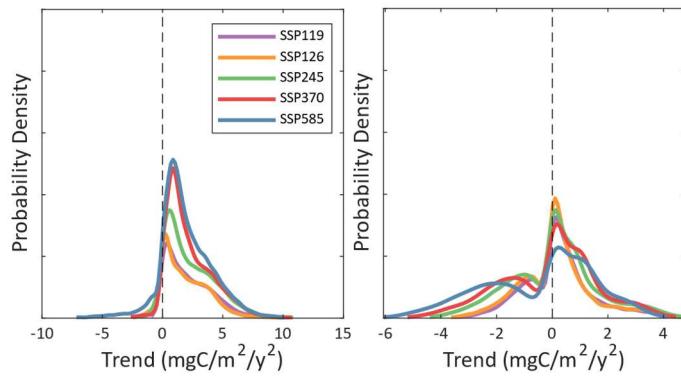
Scenario names	Shared Socioeconomic Pathways (SSPs)	Climate forcing levels	Included general circulation model (GCMs)
SSP585	SSP5	RCP8.5	CanESM5, MRI-ESM2-0, GFDL-ESM4
SSP370	SSP3	RCP7.0	CanESM5, MRI-ESM2-0, GFDL-ESM4
SSP245	SSP2	RCP4.5	CanESM5, MRI-ESM2-0, GFDL-ESM4
SSP126	SSP1	RCP2.6	CanESM5, MRI-ESM2-0, GFDL-ESM4
SSP119	SSP1	RCP1.9	CanESM5, MRI-ESM2-0, GFDL-ESM4



68
69 *Extended Data Figure 4 Anomalies of annual mean temperature (a), precipitation (b) and surface shortwave radiation (c)*
70 *over 1985-2100. The period 1985-2014 has been used as the base line to calculate anomalies.*

71 Total isoprene and monoterpene emissions

72 The distribution of significant trends in isoprene emissions reveals that a larger area displays a positive trend in
73 isoprene emission in predictions with greater levels of climate change, which is however not the same for
74 monoterpene emissions (Extended Data Fig. 5). For monoterpene, the higher CO₂ emission scenarios (such as
75 SSP585) together with warmer climate conditions result in an increased area with significant negative trends,
76 and a decrease in area with positive trends.

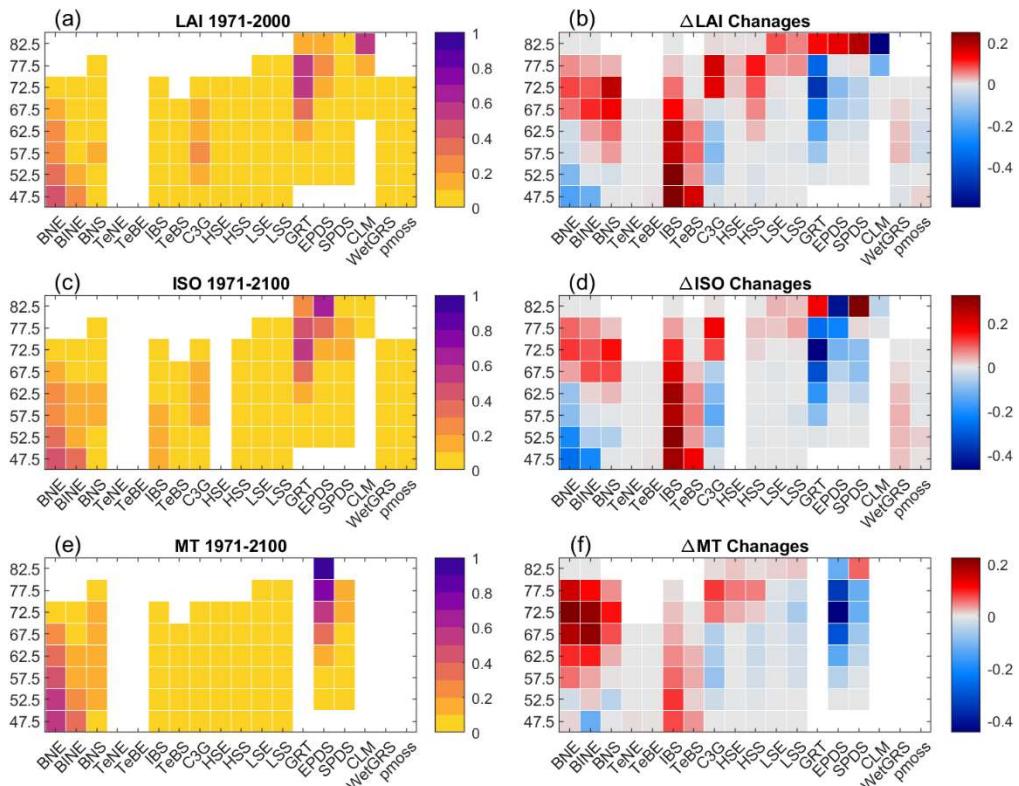


77

78 *Extended Data Figure 5. Probability density of annual isoprene (left) and monoterpene (right) emission trends. The trends*
 79 *were calculated based on the averaged emissions over 3 GCMs under each SSP, and only significant trends (Mann-Kendall*
 80 *trend test, $p < 0.05$) are shown and included in plot.*

81

82 The modelled latitudinal fractions of LAI, isoprene and monoterpene emissions for each PFT are shown in
 83 Extended Data Fig. 6. During the historical period 1971-2000, there are no emerging temperate evergreen
 84 species (represented by PFTs TeBE and TeNE in this region), but these two PFTs has started to appear in the
 85 future period 2071-2100 in the CanESM5 SSP585 scenario. The modelled changes in LAI, annual isoprene and
 86 monoterpene emissions (Extended data Fig. 6 b, d, f) show that: (1) in the high Arctic (north of 70 °N), there are
 87 large increases in the abundance of different heights of shrubs, cold grass (C3G) as well as boreal needle-leaved
 88 trees. The isoprene emissions increases are mainly contributed by the emissions from GRT, SPDS, as well as
 89 different boreal tree PFTs. The increases of monoterpene emissions in this region result mainly from emissions
 90 from boreal needle-leaved tree PFTs. (2) in the low Arctic and boreal region, there are widespread increases of
 91 IBS and TeBS (mainly in replacing of BNE and BINE, see Extended Data Fig. 6b), which result in a large
 92 increase of isoprene emissions from these two new, dominant PFTs. The increase of these two PFTs also
 93 contributes to a slight increase of monoterpene emissions, but the PFTs they replace, i.e., BNE and BINE, show
 94 large decreases, leading to a net decrease in monoterpene emission for these southern latitudinal bands.



95

96 *Extended Data Figure 6. Latitudinal fractions of leaf area index (LAI), annual isoprene (ISO) and annual monoterpane (MT)*
 97 *emissions for each modelled plant functions types (PFT, on the x-axis). The fractions of all PFTs within each latitudinal band*
 98 *add up till 1. The left column shows the modelled LAI, ISO and MT in latitudinal fractions for the period 1971-2000, and the*
 99 *right column shows the corresponding changes between 2071-2100 and 1971-2000. The future run was taken from CanESM5*
 100 *SSP585. BNE: Boreal needle-leaved evergreen; BINE: Boreal shade-intolerant needle-leaved evergreen; BNS: Boreal*
 101 *needle-leaved summergreen; TeNE: Temperate needle-leaved evergreen; TeBE: Temperate broad-leaved evergreen; C3G:*
 102 *Cool grass; HSE: High shrubs evergreen; HSS: High shrubs summergreen; LSE: Low shrub evergreen; LSS: Low*
 103 *shrub summergreen; GRT: Graminoid and forb tundra; EPDS: Evergreen prostrate dwarf shrub; SPDS: Summergreen prostrate*
 104 *dwarf shrub; CLM: Cushion forb, lichen and moss; WetGRS: flood-tolerant grass; pmoss: peatland moss.*

105 Factorial runs

106 Driven by climate data from CanESM5 SSP119 and CanESM5 SSP585, different factorial experiments are
 107 implemented (See Extended data Table 3). The associated effects are calculated as the differences between two
 108 runs.

109 *Extended Data Table 3 Overview of factorial experiments conducted with LPJ-GUESS following CanESM SSP119 and*
 110 *SSP585.*

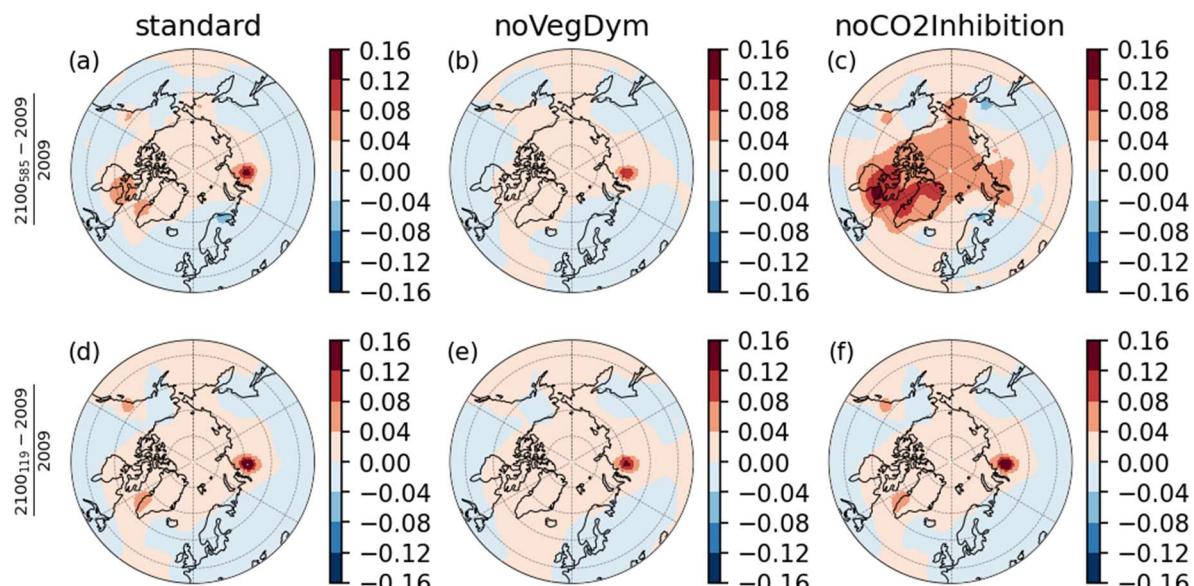
Names of different runs	CO ₂ inhibition	CO ₂ fertilization of photosynthesis	CC-induced Vegetation Changess*	Nitrogen (N) limitation	CC impacts on BVOC productions*	Effects to analyse for 2015-2100
Standard run	Yes	Yes	Yes	Yes	Yes	Full Set
noCO₂inhibition	No	Yes	Yes	Yes	Yes	CO ₂ inhibition = Standard run - noCO ₂ inhibition
noCO₂	No	No	Yes	Yes	Yes	CO ₂ fertilization = noCO ₂ inhibition - noCO ₂
noNlim	Yes	Yes	Yes	No	Yes	N limitation = Standard runs - noNlim

noVegDym	Yes	Yes	No	Yes	Yes	Vegetation changes = Standard run- noVegDym
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111 *CC: Climate change

112 TM5 modelled cloud condensation nuclei

113 The modelled cloud condensation nuclei (CCN) at 1.0% supersaturation from standard, noVegDym and
 114 noCO2Inhibition runs are further compared below. CCN (1.0%) roughly represents the number concentration of
 115 particles larger than 50 nm in diameter, so it is sensitive to new particle formation and growth which are affected
 116 by the gas precursors ELVOCs (extreme low volatile organic compounds) and SVOCs (semi-volatile organic
 117 compounds)¹⁴. Extended Data Figure 7 shows that a northward shift of CCN (1.0%), indicating potential more
 118 clouds at high latitude and less at mid latitude. This could result in global warming effect since less shortwave
 119 radiation is reflected by the same amount of clouds at high latitude compared to that at mid latitude, which is
 120 similar with the mid-latitude cloud reflectance feedback (Fig. 6 in ¹⁵).



121

122 *Extended Data Figure 7 CCN concentration at 1.0% supersaturation in the unit of [#/cm⁻³] and two factorial experiments,*
 123 *namely noVegDym and noCO2Inhibition. The top and low panels show the results driven by CanESM5 SSP585 and*
 124 *CanESM5 SSP119, respectively.*

125

126 Model uncertainties

127 *Uncertainties in modelled vegetation changes*

128 The northward shifts of woody plants as well as changes to PFT compositions simulated by LPJ-GUESS
 129 consider PFT competition and PFT responses to changing climatic and environmental conditions, including soil
 130 conditions and nutrients availability. Migration and establishment rates may be overestimated as constraints such

131 as seed dispersal have not been accounted for in the model¹⁶. However, tree demography and competition rather
132 than seed dispersal have been shown to be more important in limiting vegetation shifts in the Alps¹⁷. Our
133 modelled vegetation responses to climate (e.g., replacement of evergreen trees with deciduous trees,
134 shrubification and northward movements of evergreen trees) are consistent with experimental evidences¹⁸ and
135 other modelling studies^{4,19}.

136 *Uncertainties in modelled aerosol changes*

137 Other uncertainties of the TM5 model itself also applied to this study. For example, the wet removal of aerosol
138 particles may be overestimated, which results in lower aerosol optical depth (AOD) values^{20,21}. However, the
139 AOD comparisons of two different runs with the absolute differences can help to offset the overestimation, and
140 the relative differences may be overestimated due to underestimated AOD values. The considered SOA
141 formation process can also contribute to the uncertainties, partly due to the complicated mechanism and partly
142 due to the relatively simplified implementation in the large-scale model, as in TM5.

143

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